

Specific Capacity = discharge rate/max drawdown
 after pumping at a constant, known rate for a time until apparent equilibrium is reached (i.e., minimal change in drawdown with time)

Theis et al, 1963 - Theory - USGS Water Supply Paper 1536
 Bradbury & Rothschild, 1985 - Computer Application - Ground Water, v23, n2 240-245.
 Czarnecki & Craig, 1985 - Hand Calculator Application - Ground Water, v23, n5, p. 667-672.

Rearrange Theis (Jacob simplification) Equation to estimate T from Specific Capacity:

$$T = \frac{Q}{4\pi s} \left(-0.5772 - \ln\left(\frac{r^2 S}{4T}\right) \right)$$

must assume an S
 Note T on both sides

$$f(T) = T - \frac{Q}{4\pi s} \left(-0.5772 - \ln\left(\frac{r^2 S}{4T}\right) \right)$$

Rearrange and find a T value such that f(T) approaches zero
 assume an S
 correct for well losses ($s=s_w$)
 correct for partial penetration (Bradbury)

***** IN THE PUMPING WELL**

What is specific capacity?

Where do we get specific capacity?

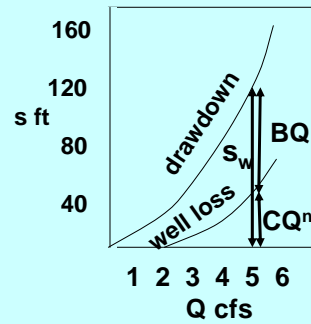
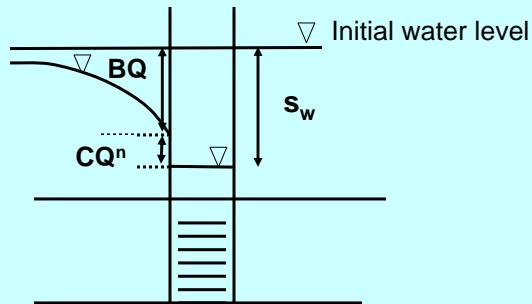
Why do we want to calculate Transmissivity from specific capacity?

Beware - using water level in **pumped well** in analyses
Beware - using well hydraulics equation's to predict w.l. in **pumped well** for design purposes

Drawdown is partially due to flow through aquifer, but head loss is also caused by flow through screen and well bore to the pump

Energy is dissipated in turbulent flow and reflected as lower head in well
 Magnitude of well loss depends on discharge velocity, minimize loss by keeping V low at Steady State described as:

$$s_{well} = \frac{Q}{2\pi T} \ln \frac{r_2}{r_1} + CQ^n$$

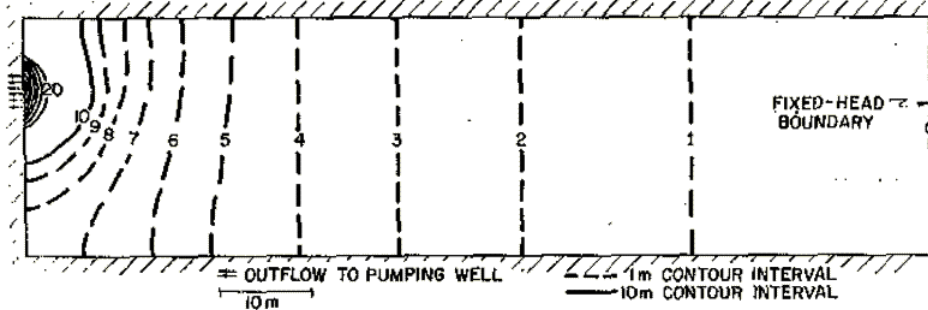


Partial Penetration

Full Penetration Equations are OK for:

$$r > 1.5b\sqrt{K_H/K_V}$$

beyond this r the equipotential lines are vertical and equivalent to the values that would be obtained from a full penetrating well

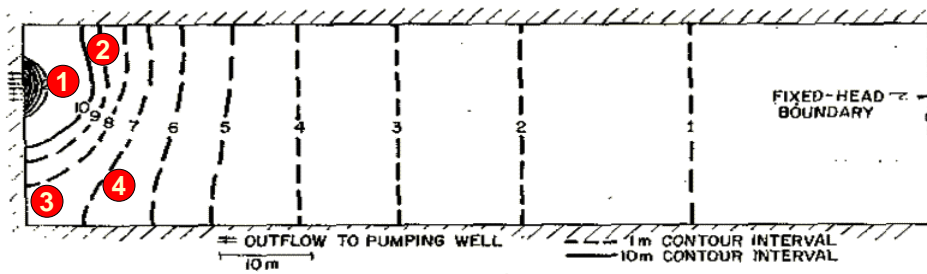




If $K_H = 4$ $K_V = 1$ $b = 200$ ft

How far must I be from the well to avoid affects of partial penetration?

$$r > 1.5b\sqrt{K_H/K_V}$$



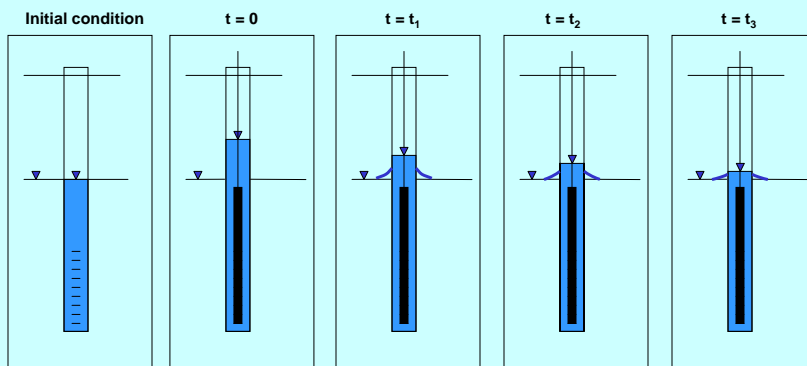
Slug Testing

Slug tests are conducted by "instantaneously" raising or lowering the water level in a well and monitoring the recovery of the water level

Often accomplished by dropping a long object into the well to displace the water

Preferable to adding a slug of water to the well which influences the chemistry

Sometimes a slug is removed, or water is bailed, from the well, decreasing the water level



Slug Testing

Cooper - Bredehoeft - Papadopoulos

- offer a solution for a slug test in a confined aquifer via curve matching
- BEWARE! Storage coefficient estimated from this approach is not reliable
- we will not go into these in this class ... curve matching, same as before

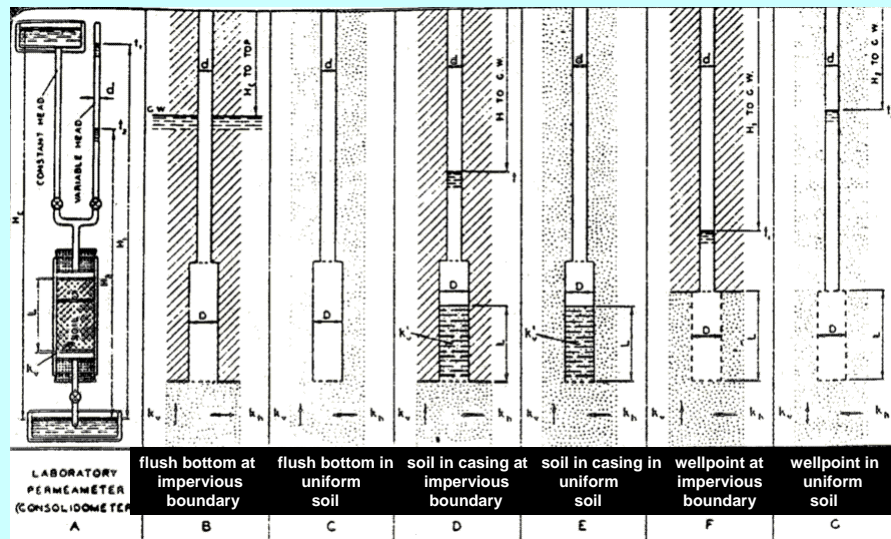
Hvorslev

- assumes water level change in the aquifer can be ignored
- offered many solutions for both confined and unconfined, see his:
Waterway Experiment Station - Army Corps of Engineers
Bulletin No. 36 - April 1951
Time Lag and Soil Permeability

Bouwer & Rice

- assume water level change in the aquifer can be ignored with the exception of its affect on geometry via the effective radius of influence
- also offer a solution for unconfined aquifers using some empirically developed coefficients

Hvorslev developed relationships for many configurations
a few examples:



As an example of one of Hvorslev's relationships, take $L_e / R > 8$:

plot h/h_0 vs. time on semi-log paper,
the slope is $\ln(h_1/h_2) / (t_1 - t_2)$

then for consistent units where:

r - casing radius

L_e - effective well screen length

R - effective well screen radius

$$K = \frac{r^2 \ln\left(\frac{L_e}{R}\right) \ln\left(\frac{h_1}{h_2}\right)}{2 L_e (t_2 - t_1)}$$

We can simplify to use $T_0 = (t_1 - t_2) =$
time to reach 37% remaining to recover

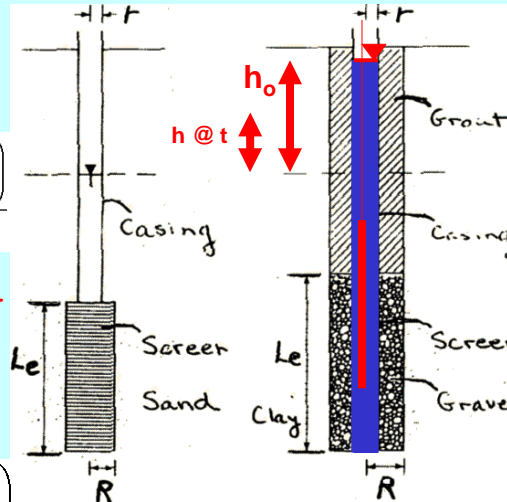
take $h_1 = h_0$ @ $t_1 = 0$

and $h_2 = 0.37 h_0$ @ $t_2 =$ when 37%
remains to recover

then $\ln(h_1/h_2) = \ln(h_0 / 0.37h_0)$
 $= \ln 2.7 = 1.0$

so slope is $1/T_0$

$$K = \frac{r^2 \ln\left(\frac{L_e}{R}\right)}{2 L_e T_0}$$



Bouwer and Rice offer:

$$K = \frac{r_c^2 \ln\left(\frac{R_e}{R}\right)}{2 L_e} \frac{1}{t} \ln\left(\frac{H_0}{H_t}\right)$$

Hvorslev

$$K = \frac{r^2 \ln\left(\frac{L_e}{R}\right) \ln\left(\frac{h_1}{h_2}\right)}{2 L_e (t_2 - t_1)}$$

rate of change of h with time

think of $\frac{1}{t} \ln \frac{H_0}{H_t}$ as $\frac{1}{t_2 - t_1} \ln \frac{H_1}{H_2}$

consistent units

r_c - casing radius

R - effective well screen radius

R_e - effective radius of head dissipation

L_e - effective well screen length

H_0 - drawdown at time = 0

H_t - drawdown at time = t

R_e is difficult to determine

$$K = \frac{r_c^2 \ln\left(\frac{R_e}{R}\right)}{2 L_e} \frac{1}{t} \ln\left(\frac{H_o}{H_i}\right)$$

R_e is difficult to determine

Bower and Rice undertook lab experiments in sand tanks to establish $\ln(R_e/R)$

for $L_w < h \ln\left(\frac{R_e}{R}\right) = \left[\frac{1.1}{\ln\left(\frac{L_w}{R}\right)} + \frac{A + B \ln\left(\frac{h-L_w}{R}\right)}{\frac{L_e}{R}} \right]^{-1}$

for $L_w = h \ln\left(\frac{R_e}{R}\right) = \left[\frac{1.1}{\ln\left(\frac{L_w}{R}\right)} + \frac{C}{R} \right]^{-1}$

Gather data from the sand tank

Work with a partner to estimate K via Hvorslev AND Bower and Rice Methods

How do the K values compare? And for the pump test we did earlier in the semester?

How does the tank fit the assumptions of the methods?

$$K = \frac{r^2 \ln\left(\frac{L_e}{R}\right)}{2 L_e T_o}$$

$$K = \frac{r_c^2 \ln\left(\frac{R_e}{R}\right)}{2 L_e} \frac{1}{t} \ln\left(\frac{H_o}{H_i}\right)$$

for $L_w < h \ln\left(\frac{R_e}{R}\right) = \left[\frac{1.1}{\ln\left(\frac{L_w}{R}\right)} + \frac{A + B \ln\left(\frac{h-L_w}{R}\right)}{\frac{L_e}{R}} \right]^{-1}$

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